Asian Journal of Probability and Statistics

Volume 25, Issue 4, Page 15-26, 2023; Article no.AJPAS.110014 *ISSN: 2582-0230*

Non-Stationary Modeling of Annual Flood Peak Heights of Mahanadi River Basin with the q-Generalized Extreme Value Distribution

S. Nagesh a++* and Laxmi. B. Dharmannavar b#

^a Department of Statistics, Karnatak University Dharwad, Karnataka-580003, India. ^b CHRIST (Deemed to be University), Yeshwanthpur Campus, Bangalore, Karnataka-560073, India.

Authors' contributions

This work was carried out in collaboration between both authors. Both authors read and approved the final manuscript.

Article Information

DOI: 10.9734/AJPAS/2023/v25i4569

Open Peer Review History:

This journal follows the Advanced Open Peer Review policy. Identity of the Reviewers, Editor(s) and additional Reviewers, peer review comments, different versions of the manuscript, comments of the editors, etc are available here: https://www.sdiarticle5.com/review-history/110014

Original Research Article

Received: 03/10/2023 Accepted: 09/12/2023 Published: 15/12/2023

Abstract

In recent years, due to climate change, catastrophic events are increased largely in India. Hence researchers are forced to consider non-stationary flood frequency analysis as an improved method. In this paper, nonstationarity of annual daily maximum flood heights were studied at 12 sites of Mahanadi River Basin (MRB) by analyzing the flood frequency of a stationary model and 4 non-stationary models using time dependent q-GEV model by considering trend as a linear function of its location and scale parameters. The q-GEV distribution is utilized in this study because of its flexibility and accuracy than GEV distribution in modeling extreme flood heights. The results found that there is strong evidence of a linear trend existence for both the location and scale parameters at the Kesinga site; for the location parameter at Pathardi and Simga sites; for the scale parameter at

__

__

Asian J. Prob. Stat., vol. 25, no. 4, pp. 15-26, 2023

⁺⁺Associate Professor;

[#] Assistant Professor;

^{}Corresponding author: Email: nageshstatistics@gmail.com;*

Dharmajagarh, Kotni and Seorinarayan, and no linear trend exists for both location and scale parameters at Alipingal, Bomnidhi, Manendragarh, Mohana, Rajim and Sundargarh, there may be exists other form of trend at these sites. The findings also indicate that nonstationarity is present in the MRB due to climate change, which help to water practitioner for taking precautions against adverse effect of extreme floods.

Keywords: Mahanadi river basin; non-stationarity; q-GEV model; MLE.

1 Introduction

In the past few years, occurrence of extreme hydrological events such as floods, droughts are increasingly repeated due to impact of climate change, urbanization, deforestation and encroachment of river basin, particularly in India [1,2]. In India, each year on an average 1600 persons die as a result of the floods,about Rs. 5600 Crores (73 Million USD) as a damage cost due to floods and 12% of geographic area of India is flood prone [2]. The assumption of stationary is that in a system, statistical characteristics such as mean and variance do not vary over time and show no trend [3]. The impact of climate change, global warming, deforestation and urbanization can invalidate the assumption of stationarity [4], Gumbel [5] expressed invalidation of stationary model flood frequency analysis in the situations of climate change and other variations.It is therefore necessary to develop alternative methods that take nonstationarity into account for the effective design of hydraulic control structures. The floods are exacerbated by natural events such as reduced carrying capacity of the water course caused by silting of the river bed, landslide causing obstacles and change in river path and by manmade events such as unplanned urbanization and floodplain encroachment [6]. The flood frequency analysis of one stationary and 12 non-stationary models were used to evaluate the nonstationarity of Periyar River flow, and their results indicate that climate change and anthropogenic activities are equally responsible for the nonstationarity of the Periyar River discharge [6]. Seasonality and Trend in natural occurrences such as floods are common causes of nonstationarity[7]. In non-stationary settings, Hounkpe[8] provided a statistical model for predicting flood probability, the model was used to five gauging stations in the Queme River Basin in Benin Republic, West Africa and it fits the annual maximum discharge using a time-dependent and covariate-dependent generalized extreme value (GEV) distribution.

In the above works, the GEV distribution was considered to be a good model for non-stationary modeling with different approaches. Nagesh and Laxmi [9]: identified GEV distribution as a best flood frequency distribution for 12 sitesin MRB under the stationary conditions. In this study, extension of Nagesh and Laxmi [10] work, by utilizing an extended GEV model q-GEV distribution for the first time as an alternate distribution for modeling the non-stationary situation in MRB. Nagesh and Laxmi [10]: Estimating the Extreme Flood Height Quantiles through Bayesian Approach using q-GEV distribution in MRB.

In practice, the GEV model may be insufficient, but its generalizations should provide more modelling flexibility. Provost et al. [11] introduced the extended model q-GEV distribution, which is a q analogue of the GEV, where q is an extra parameter that allows for greater flexibility in modeling extreme events than the GEV distribution. In previous work, the q-GEV distribution was shown that better flood frequency model than GEV distribution for twelve MRB sites using the block maxima technique with stationary assumption. For the same sites, our objective in this study is to model maximum flood height using time dependent q-GEV distribution under non-stationary assumption that is when covariates are present. This study involves a non-stationary q-GEV distribution as time dependent, expecting the variation linearly or nonlinearly over time in location and scale parameters while shape parameter is unchanged with the time. A statistical modeling approach is advocated by considering maximum likelihood estimation (MLE) in the presence of covariates like trends and cycles [12]. Further related information can be found in the literature [13-15].

As far as we know no similar work has been done in earlier studies on extreme value theory in a changing climate for the Mahanadi River Basin.

2 Materials and Methods

This section explains how the data was analyzed. The extended time homogeneous q-GEV distribution was used to investigate linear and quadratic trend models.

2.1 The data

The original form of twelve sites flood height data recorded thrice a day was provided by the Central Water Commission (CWC), Bhubaneshwar, which is the competent authority for water resources management in the river basin. Using successive steps in each hydrological year, the annual maximum series was obtained. Ferreira and De Haan [16] provide detailed descriptions of the block maxima probability theory as well as practical considerations for choosing block maxima rather than peak over threshold, Dombry[17] demonstrated that when utilizing the block maxima technique ML estimators are consistent. Block maxima and Peak over Threshold are two important methodologies used in extreme value theory for flood frequency analysis. When the sample size is large, the block maxima approach is utilized, in which each hydrological year is considered as its own block [16].

2.2 Non-stationary extreme value models

The goal of the study was to model the behavior of annual maximum flood heights in the presence of covariates, in order to see if non-stationary models fit the data better than stationary (time independent) models. Trends, cycles, and physical factors are examples of covariates, according to Katz et al. [12]. Trends are considered as covariates in particular, and the method of time varying moments is the most commonly employed approach to non-stationary flood frequency analysis.

The models M_0 , M_1 , M_2 , M_3 , and M_4 were considered to compare the stationary (Time independent) and nonstationary (Time dependent) models, where

 $M₀$ - q-GEV time independent (stationary) model, M_1 - q-GEV model with linear trend in location and scale parameter, M_2 - q-GEV model with linear trend in location parameter, M3 - q-GEV model with linear trend in scale parameter and M4 - q-GEV model with non-linear trend in location parameter.

In previous work, the best flood frequency model was shown to be the q-GEV (time independent) distribution at twelve MRB sites [9]. This study looks at non-stationary (time-dependent) models for the same sites. M_0 is a stationary model that is not affected by time. Non-stationary models M_1, M_2, M_3 , and M_4 have a linear or nonlinear trend in the location or scale parameter, or both.

The distribution function and density function of q-GEV distribution is given by

$$
F(x; s, m, \xi, q) = \left[1 + q\left(1 + \xi(xs - m)\right)^{-\frac{1}{\xi}}\right]^{-\frac{1}{q}}; \xi \neq 0, q \neq 0
$$
\n(1)

$$
f(x; s, m, \xi, q) = s\left(1 + \xi(xs - m)\right)^{\left(-\frac{1}{\xi}\right) - 1} \left[1 + q\left(1 + \xi(xs - m)\right)^{\left(-\frac{1}{\xi}\right)}\right]^{\left(-\frac{1}{q}\right) - 1}; \ \xi \neq 0, q \neq 0 \tag{2}
$$

The log-likelihood function of q-GEV distribution is given by

$$
l(s, m, \xi, q) = n * \log(s) + \left(-\frac{1}{q} - 1\right) \sum_{i=1}^{n} \log\left[q(\xi(x_i s - m) + 1)^{-1/\xi} + 1\right] + \left(-\frac{1}{\xi} - 1\right) \sum_{i=1}^{n} \log\left[\xi(x_i s - m) + 1\right]
$$
\n(3)

where $m = \mu/\sigma$ and $s = 1/\sigma$, μ is location parameter, σ is scale parameter, ξ and η are shape parameters. The distribution function, Probability density function and log-likelihood function of q-GEV distribution without re-parameterization are given by equations (4), (5) and (6) respectively.

$$
F(x; \mu, \sigma, \xi, q) = \left\{ 1 + q \left[1 + \xi \left(\frac{x - \mu}{\sigma} \right) \right]^{-1/\xi} \right\}^{-1/q}; \xi \neq 0, q \neq 0 \tag{4}
$$

$$
f(x; \mu, \sigma, \xi, q) = \frac{1}{\sigma} \Big[1 + \xi \Big(\frac{x - \mu}{\sigma} \Big) \Big|^{(-\frac{1}{\xi}) - 1} \left\{ 1 + q \Big[1 + \xi \Big(\frac{x - \mu}{\sigma} \Big) \Big]^{-\frac{1}{\xi}} \right\}^{(-\frac{1}{q}) - 1}; \ \mu \in R, \sigma > 0, \xi \neq 0, q \neq 0 \tag{5}
$$

17

$$
l(\mu, \sigma, \xi, q) = n * \log(\sigma) + \left(-\frac{1}{\xi} - 1\right) \sum_{i=1}^{n} \log\left[1 + \xi\left(\frac{x_i - \mu}{\sigma}\right)\right] + \left(-\frac{1}{q} - 1\right) \sum_{i=1}^{n} \log\left\{1 + q\left[\xi\left(\frac{x_i - \mu}{\sigma}\right)\right]^{-1/\xi}\right\} \tag{6}
$$

M₁ is non-stationary q-GEV model with linear trend in both location and scale parameter is $\mu(t) = \mu_0 +$ $\mu_1 t$, $\log \sigma(t) = \sigma_0 + \sigma_1 t$, $\xi(t) = \xi$ and $q(t) = q$

The distribution function of M_1 is given by equation (7) while the log-likelihood function of M_1 is given by equation (8).

$$
F(x; \mu(t), \sigma(t), \xi(t), q(t)) = \left\{ 1 + q \left[1 + \xi \left(\frac{x - (\mu_0 + \mu_1 t)}{\exp(\sigma_0 + \sigma_1 t)} \right) \right]^{-1/\xi} \right\}^{-1/q}; \xi \neq 0, q \neq 0
$$
\n
$$
l(\mu(t), \sigma(t), \xi(t), q(t)) = n * \log \left(-\frac{1}{\xi} - 1 \right) \sum_{i=1}^{n} \log \left[1 + \xi \left(\frac{x_i - (\mu_0 + \mu_1 t)}{\exp(\sigma_0 + \sigma_1 t)} \right) \right] + \left. \right.
$$
\n
$$
= \left(-\frac{1}{q} - 1 \right) \sum_{i=1}^{n} \log \left\{ 1 + q \left[\xi \left(\frac{x_i - (\mu_0 + \mu_1 t)}{\exp(\sigma_0 + \sigma_1 t)} \right) \right]^{-1/\xi} \right\}
$$
\n
$$
(8)
$$

where t is time (in years). The log-likelihood function of M_1 is given by equation (8), and the set of values is estimated by maximization of the log-likelihood function. The Newton Raphson method is used to solve the loglikelihood function equations.

M² is non-stationary (time dependent) q-GEV model with linear trend in the location parameter.

$$
\mu(t) = \mu_0 + \mu_1 t, \sigma(t) = \sigma, \xi(t) = \xi \text{ and } q(t) = q
$$

Cumulative distribution function and log-likelihood function of M_2 is given by equation (9) and equation (10) respectively

$$
F(x; \mu(t), \sigma(t), \xi(t), q(t)) = \left\{ 1 + q \left[1 + \xi \left(\frac{x - (\mu_0 + \mu_1 t)}{\sigma} \right) \right]^{-1/\xi} \right\}^{-1/q}; \ \xi \neq 0, q \neq 0(9)
$$

$$
l(\mu(t), \sigma(t), \xi(t), q(t)) = n * \log(\sigma) + \left(-\frac{1}{\xi} - 1 \right) \sum_{i=1}^{n} \log \left[1 + \xi \left(\frac{x_i - (\mu_0 + \mu_1 t)}{\sigma} \right) \right] +
$$

$$
\left(-\frac{1}{q} - 1 \right) \sum_{i=1}^{n} \log \left\{ 1 + q \left[\xi \left(\frac{x_i - (\mu_0 + \mu_1 t)}{\sigma} \right) \right]^{-1/\xi} \right\}
$$
 (10)

M³ is non-stationary q-GEV model with linear trend in the scale parameter.

$$
log \sigma(t) = \sigma_0 + \sigma_1 t
$$
, $\mu(t)=\mu$, $\xi(t) = \xi$ and $q(t) = q$

Distribution function and log-likelihood function of M_3 model are obtained by replacing above quantities in equations (4) and (6) respectively.

$$
F(x; \mu(t), \sigma(t), \xi(t), q(t)) = \left\{ 1 + q \left[1 + \xi \left(\frac{x - \mu}{\exp{(\sigma_0 + \sigma_1 t)}} \right) \right]^{-1/\xi} \right\}^{-1/q}; \xi \neq 0, q \neq 0 \tag{11}
$$

$$
l(\mu(t), \sigma(t), \xi(t), q(t)) = n * \ln(\sigma_0 + \sigma_1 t) + \left(-\frac{1}{\xi} - 1\right) \sum_{i=1}^n \log\left[1 + \xi\left(\frac{x_i - \mu}{\exp(\sigma_0 + \sigma_1 t)}\right)\right] +
$$

18

Nagesh and Dharmannavar; Asian J. Prob. Stat., vol. 25, no. 4, pp. 15-26, 2023; Article no.AJPAS.110014

$$
\left(-\frac{1}{q}-1\right)\sum_{i=1}^{n}\log\left\{1+q\left[\xi\left(\frac{x_i-\mu}{\exp\left(\sigma_0+\sigma_1t\right)}\right)\right]^{-1/\xi}\right\}\tag{12}
$$

M4 is non-stationary q-GEV model with non-linear quadratic trend in the location parameter.

 $\mu(t) = \mu_0 + \mu_1 t + \mu_2 t^2, \sigma(t) = \sigma, \xi(t) = \xi \text{ and } q(t) = q$

Similarly by replacing above quantities in equations (4) and (6) we can get distribution and log-likelihood function of M⁴ in q-GEV Model. 1

$$
F(x; \mu(t), \sigma(t), \xi(t), q(t)) = \left\{ 1 + q \left[1 + \xi \left(\frac{x - (\mu_0 + \mu_1 t + \mu_2 t^2)}{\sigma} \right) \right]^{-\frac{1}{\xi}} \right\}^{-\frac{1}{q}}; \xi \neq 0, q \neq 0 \tag{13}
$$

$$
l(\mu(t), \sigma(t), \xi(t), q(t)) = n * \log(\sigma) + \left(-\frac{1}{\xi} - 1\right) \sum_{i=1}^{n} \log\left[1 + \xi \left(\frac{x_i - (\mu_0 + \mu_1 t + \mu_2 t^2)}{\sigma}\right)\right] + \left(-\frac{1}{q} - 1\right) \sum_{i=1}^{n} \log\left\{1 + q\left[\xi \left(\frac{x_i - (\mu_0 + \mu_1 t + \mu_2 t^2)}{\sigma}\right)\right]^{-1/\xi}\right\}
$$
(14)

Using the maximum likelihood estimation technique, parameters of the models M_0 , M_1 , M_2 , M_3 , and M_4 are calculated using MLE technique. MLE is used to estimate the parameters of both stationary and non-stationary q-GEV models. Using the MLE methodology in the presence of variables is reliable in both the Block Maxima and Peaks over Threshold procedures [18,19].

2.3 Model choice

To compare one model to the other, the MLE of nested models employs a simple procedure known as the Deviance (D) statistic [18-20]. The time-homogeneous GEV model, M_0 , is a subset of the time-dependent models M_1 , M_2 , M_3 , and M_4 in this research. To check significance of non-stationary models over stationary model D statistic is given by

$$
D = 2\{l_i(M_i) - l_0(M_0)\}\tag{15}
$$

where $l_i(M_i)$ is the maximized negative log-likelihood value of i^{th} model ($i = 1,2,3,4$), $l_0(M_0)$ is the maximized negative log-likelihood value of time independent(stationary) model. D follows chi-square distribution with k degrees of freedom, where k is difference in number of parameters. If $D > \chi^2_{k,\alpha}$, then we reject H₀. It suggests that M_i is more significant than M_0 .

3 Results and Discussion

The parameters of the models M_0 , M_1 , M_2 , M_3 , and M_4 were calculated using MLE. Parameter estimates of nonstationary q-GEV model and Deviance statistics calculated for pairs of stationary and non-stationary models for Bamnidhi site are given in Table1. For Bamnidhi site, consider models M_0 and M_1 , negative log-likelihood value for M_0 and M_1 is -52.9804 and -51.4985 respectively.

Using equation (15), D=2[(-51.4985) - (-52.9804)] = 2.9639, which is less than 5.991, i.e. $\chi_{2,0.05}^2$ = 5.991. So we can conclude that M_0 is better fit than M_1 for Bamnidhi site.

Table 1. Parameter estimates of annual maximum time heterogeneous q-GEV models for Bamnidhi site

Model	μ_0	u	u_2	σ0	σ 1				
\mathbf{M}_0	4.6943			0.8055	0	-0.5750	1.7019	-52.9804	
M_1	.4083	.408		0.9667	0.4028	-0.6900	1.5317	-51.4985	2.9639
M ₂	4.2001	0.261		0.6427		-0.4263	1.2799	-51.5775	2.8058
M_3	4.8455			0.7180	-0.0287	-0.7785	1.0460	-51.9625	2.0359
M4	4.6503	.258	0.217	0.6620	0	-1.1672	2.6899	-52.6737	0.6135

In other words non-stationary model with linear trend in location parameter is insignificant. It clearly shows that the non-stationary model does not give any improvement in fit over the time-homogeneous q-GEV model. Here we can conclude that M_0 is better fit than M_1 . Next consider M_0 and M_2 .

Negative log-likelihood value for M_0 and M_2 , is -52.9804 and -51.5775 respectively.

 $D = 2[(-51.5775) - (-52.9804)] = 2.8058$, which is less than 3.8414. $\chi^{2}_{1,0.05} = 3.8414$.

Similarly the negative log-likelihood value and D statistics values for other pairs of models (M_0, M_3) and (M_0, M_4) are given in Table1. The D statistic values are less than critical values (i.e. 2.0359 < 3.8414 and 0.6135 < 5.991) for the models respectively, p-values when $\mu=0$ and $\sigma=0$ are not less than 0.05 for the above models. The models (M_1, M_2, M_3, M_4) do not provide any improvement in fit over the time-homogeneous q-GEV model at Bamnidhi site. Therefore stationary model for Bamnidhi site is given by (using equation 4).

$$
F(x; \mu, \sigma, \xi, q) = \left\{ 1 + 1.7019 \left[1 - 0.575 \left(\frac{x_i - 4.6943}{0.8055} \right) \right]^{1/0.575} \right\}^{-1/1.7019}
$$

Table 2. Parameter estimates of annual maximum time heterogeneous q-GEV models for Dharamjaigarh site

Model	LU ₀	$\mathsf{\mu}_1$	\mathbf{u}_2	σ_0	σ 1				
M_0	6.0787			0.7707		0.3404	0.6506	-44.7622	
M_1	6.0180	0.6079		0.6166	0.4624	0.3106	0.9425	-43.1313	3.2618
$\rm M_2$	5.8273	0.0562		0.3776		0.2542	0.9349	-44.2648	0.9947
M_3	6.7685			0.6982	-0.049	0.2014	0.7433	-42.6482	4.2279
$\rm M_4$	5.5320	0.2170	0.006	3.0090		0.3138	0.8911	-43.9958	.5327

Consider the models (M₀, M₁), D statistic is 3.2618 which is less than $\chi_{2,0.05}^2 = 5.991$

Linear trend in location parameter is not significant at 5% level of significance. D statistic value for the models (M_0, M_2) is 0.9947, which is less than 3.8414(from Table2). D statistic value for the models (M_0, M_3) is 4.2279, which is greater than 3.8414.

The likelihood ratio test for $\sigma_1 = 0$ has p-value 0.0148 (indicates that M_3 is significant). It is worthwhile to consider time-heterogeneous (non-stationary) model M_3 that is linear trend in scale parameter at Dharamjaigarh. So we conclude that non-stationary model (M_3) outperform than stationary model.

Therefore non-stationary model with linear trend in scale parameter for Dharamjaigarh is given by (using equation 11)

$$
F(x; \mu(t), \sigma(t), \xi(t), q(t)) = \left\{ 1 + 0.7433 \left[1 + 0.2014 \left(\frac{x_i - 6.7685}{\exp(0.6982 - 0.0489 t)} \right) \right]^{-1/0.2014} \right\}^{-1/0.7433}
$$

Parameter estimates of stationary and non-stationary q-GEV models for remaining ten sites are given in Table3.

Table 3. Parameter estimates of time heterogeneous q-GEV models for ten sites

Site name	Model	μ_0	μ1	μ	σ_0	σ_1		
Kesinga	M_0	10.0670	Ω	0	1.8674	θ	-0.6814	1.2930
	M_1	10.9203	1.0109	θ	1.8639	0.1868	-0.6156	1.1996
	$\rm M_2$	11.4792	1.0358	θ	2.6923	Ω	-0.2112	2.4827
	M_3	10.2259	Ω	0	1.954	0.6547	-0.6952	2.3249
	M4	10.0927	-4.2634	0.2132	1.9666	Ω	-0.3211	1.3031
Kotni	$\rm M_0$	9.5410	Ω	0	1.7674	Ω	-0.5996	1.2471
	\mathbf{M}_1	9.1822	0.9634	θ	1.7580	0.1770	-0.6254	1.1993
	M ₂	9.1917	1.0535	Ω	1.8565	Ω	-0.3397	1.5492
	M_{3}	8.5328	Ω	0	0.3842	-0.031	-0.7229	1.1002
	$\rm M_4$	9.9001	-0.7310	0.0416	0.5767	Ω	-0.4306	2.0120
Manendragarh	$\rm M_0$	4.3112	Ω	0	0.5755	θ	0.1531	0.2173

−1/0.7433

Table 4 gives Negative log-likelihood (NLL), D statistic values and χ^2 critical value of stationary and nonstationary time series models for remaining ten sites.

At Kesinga site, D statistic of models (M₀, M₃) is 4.8274 which is greater than tabulated value 3.8414, it shows that model M_3 is better than stationary model. The D statistic values for all other models except model (M_0 , M_3) are less than tabulated value. Therefore at Kesinga non-stationary q-GEV model with linear trend in scale parameter M_3 gives improvisation over stationary model M_0 .

So non-stationary model with linear trend in scale parameter for Kesinga site is given by

$$
F(x; \mu(t), \sigma(t), \xi(t), q(t)) = \left\{ 1 + 0.7433 \left[1 - 0.6952 \left(\frac{x_i - 10.2259}{\exp(1.954 - 0.6547t)} \right) \right]^{1/0.6952} \right\}^{-1/2.3249}
$$

Similarly from Table4, D statistic value for the pair of model M_0 and M_3 for Kotni and Seorinarayan is 4.413 and 5.6586 respectively, which are greater than tabulated value 3.8414 (i.e. $D > \chi^2_{1,0.05}$). Hence we can conclude that model M_3 is significant than M_0 for Kotni site and Seorinarayan site.

So non-stationary model with linear trend in scale parameter for Kotni site is given by

$$
F(x; \mu(t), \sigma(t), \xi(t), q(t)) = \left\{1 + 1.1002 \left[1 - 0.7229 \left(\frac{x_i - 8.5328}{\exp(0.3842 - 0.031t)}\right)\right]^{1/0.7229}\right\}^{-1/1.1002}
$$

Site name	Model	NLL	D	χ^2 Critical Value
Kesinga	\mathbf{M}_0	-83.9958		
	\mathbf{M}_1	-83.1	1.7916	5.991
	M_2	-83.3512	1.2892	3.8414
	M_3	-81.5821	4.8274	3.8414
	M_4	-83.2325	1.5266	5.991
Kotni	M_0	-82.4308		
	M_1	-82.1537	0.5542	5.991
	M ₂	-82.3325	0.1966	3.8414
	M_3	-80.2243	4.413	3.8414
	M_4	-81.9637	0.9342	5.991
Manendragarh	M_0	-44.2274		
	M_1	-43.5480	1.35876	5.991
	M_2	-43.9629	0.52896	3.8414
	M_3	-43.649	1.15676	3.8414
	\mathbf{M}_4	-43.3618	1.73116	5.991
Mohana	\mathbf{M}_0	-39.6943		
	M_1	-39.547	0.2946	5.991
	M_2	-38.9243	1.54	3.8414
	M_3	-39.254	0.8806	3.8414
	M_4	-39.1657	1.0572	5.991
Pathardhi	M_0	-51.1474		
	M_1	-50.945	0.4047	5.991
	M ₂	-48.3295	5.6357	3.8414
	M_3	-50.8495	0.5957	3.8414
	M_4	-50.5798	1.1351	5.991
Rajim	M_0	-79.2463		
	M_1	-78.5139	1.4648	5.991
	M_2	-79.1537	0.1852	3.8414
	M_3	-78.9381	0.6164	3.8414
	M_4	-78.349	1.7946	5.991
Seorinarayan	M_0	-66.772		
	M_1	-66.2973	0.9494	5.991
	M_2	-66.4967	0.5506	3.8414
	M_3	-63.9427	5.6586	3.8414
	M_4	-65.5468	2.4504	5.991
Simga	M_0	-100.255		
	M_1	-99.5826	1.345	5.991
	M_2	-98.2431	4.024	3.8414
	M_3	-99.5488	1.4126	3.8414
	M_4	-99.2849	1.9404	5.991
Sundaragarh	M_0	-54.8368		
	M_1	-53.9144	1.8448	5.991
	M_2	-54.015	1.6436	3.8414
	M_3	-54.3521	0.9694	3.8414
	M_4	-53.9986	1.6764	5.991
Alipingal	\mathbf{M}_0	-66.6641		
	M_1	-66.315	0.6981	5.991
	M ₂	-66.4983	0.3315	3.8414
	M_3	-66.3219	0.68424	3.8414
	M_4	-66.1999	0.92832	5.991

Table 4. Annual maximum time heterogeneous q-GEV models for ten sites

−1/1.9547

Non-stationary model with linear trend in scale parameter for Seorinarayan site is given by

$$
F(x; \mu(t), \sigma(t), \xi(t), q(t)) = \left\{ 1 + 1.9547 \left[1 - 1.441 \left(\frac{x_i - 13.9995}{\exp(2.0956 + 0.2354t)} \right) \right]^{1/1.441} \right\}^{-1/1.9547}
$$

Next, the D statistic values for the pair of models (M_0, M_2) of Pathardhi and Simga sites are 5.6357 and 4.024 respectively. D statistic values are greater than 3.8414 (i.e. $D > \chi^2_{1,0.05}$). Hence we can conclude that non-stationary q-GEV model with linear trend in location parameter is better model than stationary model at these sites.

Non-stationary model with linear trend in location parameter for Pathardhi is given by (using equation 9)

$$
F(x; \mu(t), \sigma(t), \xi(t), q(t)) = \left\{1 + 1.9535\left[1 - 0.6028\left(\frac{x - (6.2766 + 0.7046t)}{0.5444}\right)\right]^{1/0.6028}\right\}^{-1/1.9535}
$$

Non-stationary model with linear trend in location parameter for Simga is given by (using equation 9)

$$
F(x; \mu(t), \sigma(t), \xi(t), q(t)) = \left\{1 + 1.8502\left[1 - 0.257\left(\frac{x - (10.4293 + 0.1565t)}{0.9750}\right)\right]^{1/0.257}\right\}^{-1/1.8502}
$$

D statistics value for all the pairs of models i.e. (M_0, M_1) , (M_0, M_2) and (M_0, M_3) , and (M_0, M_4) are less than tabulated values (5.991, 3.8414, 3.8414 and 5.991 respectively) for Manendragarh, Mohana, Rajim, Sundaragrh and Alipingal. Therefore we can conclude that at these sites stationary q-GEV model is better than non-stationary models.

Stationary q-GEV model for Manendragarh is given by

$$
F(x; \mu, \sigma, \xi, q) = \left\{ 1 + 0.2173 \left[1 + 0.1531 \left(\frac{x - 4.3112}{0.5755} \right) \right]^{-1/0.1531} \right\}^{-1/0.2173}
$$

Similarly Stationary q-GEV model for Mohana, Rajim, Sundaragrh and Alipingal are obtained by putting values of μ , σ , ξ and q in the equation (4) from Table3.

D statistic value for M_0 and M_4 values of all the sites are less than 5.991, which indicates that M_4 is nowhere best model among twelve sites. It is better to ignore M_4 at all the sites. Fig. 1 gives return level plot for non-stationary model (i.e. model with linear trend in scale parameter) of Dharamjaigarh, which have higher return levels as compared to stationary model for a different return period. Same nature is found in the other non-stationary model return level plot. The return flood levels at different sites using non-stationary models are higher as compare to that of stationary models.

Fig. 1 – Fig. 6 depicts Return level plots for the sites of MRB where non-stationary q-GEV model has improvisation over stationary model.

Fig. 1. Return level plot for Dharamjaigarh site Fig. 2. Return level plot for Kesinga site

Fig. 3. Return level plot for Kotni site Fig. 4. Return level plot for Pathardhi site

Fig. 5. Return level plot for Seorinarayan site Fig. 6. Return level plot for Simga site

4 Conclusions

The study looked at the use of extreme value theory in a changing climate for the Mahanadi River Basin in India. The study looked at twelve hydrometric stations that represented twelve different sites along the MRB. The parameters of the q-GEV distribution were determined using the maximum likelihood estimation approach when a long-term trend covariate was present. The study revealed the importance of incorporating non-stationary linear and nonlinear trend models when using extreme value theory in a changing climate, as these models provide a significantly better fit than time-homogeneous models. This improvement in fitness is crucial for the government in planning and policy making.

The prevailing models at the twelve sites were successfully identified, six sites have a time-homogeneous GEV model, four sites have a prevailing time-heterogeneous GEV model with a dominant linear trend in the scale parameter, and two sites have a prevailing time-heterogeneous GEV model with a dominating linear trend in the scale parameter, according to this study.

The results of the study demonstrated that the time independent q-GEV model $(M₀)$ is much better fit than time heterogeneous q-GEV models at six sites: Alipingal, Bamnidhi, Manendragarh, Mohana, Rajim, and Sundargarh.

At sites Dharamjaigarh, Kesinga, Kotni and Seorinarayan, the time heterogeneous q-GEV model M_3 explains significant variability than M₀. Non-stationary model with linear trend in scale parameter outperformed than stationary models. At sites Pathardhi and Simga, time heterogeneous q -GEV model $M₂$ suits well as compare to M_0 . Non-stationary model with linear trend in location parameter, M_2 , is significant as compared to the Stationary model M_0 .

Non-stationary model with Non-linear trend in location parameter does not fit well over stationary models at any site, which suggests that better to drop out this model consideration for application in the MRB.

Acknowledgements

The authors thanked Central Water Commission (CWC), Bhuvaneshwar, the authority for water resource management in India under the Ministry of Jal Shakti for providing hydrometric data used in this study.

Competing Interests

Authors have declared that no competing interests exist.

References

- [1] SinghN, Chinnasamy P. Non-stationary flood frequency analysis and attribution of streamflow series: a case study of Periyar River, India. Hydrological Sciences Journal. 2021;66(13):1866-1881.
- [2] Black KP, Baba M, Mathew J, Chandra S, Singh SS, Shankar R, Kurian NP, Urich P, Narayan B, Stanley DO, ParsonsS, Ray G.Climate change adaptation guidelines for coastal protection and management in India. prepared by FCGANZDEC (New Zealand) for the Global Environment Facility and Asian Development Bank. 2018;1.
- [3] Sivakumar B. Chaos in hydrology: bridging determinism and stochasticity. Springer; 2016.
- [4] Ray LK, GoelNK, Arora M.Trend analysis and change point detection of temperature over parts of India. Theoretical and Applied Climatology. 2019;138(1):153-167.
- [5] Gumbel EJ.The return period of flood flows. The Annals of Mathematical Statistics. 1941;12(2):163-190.
- [6] Singh O, Kumar M.Flood events, fatalities and damages in India from 1978 to 2006. Natural Hazards. 2013;69(3):1815-1834.
- [7] Maposa D, Lesaoana M, Cochran JJ.Modelling non-stationary annual maximum flood heights in the lower Limpopo River basin of Mozambique. Jàmbá: Journal of Disaster Risk Studies. 2016;8(1):1-9.
- [8] HounkpèJ, Diekkrüger B, Badou DF, AfoudaAA.Non-stationary flood frequency analysis in the Ouémé River Basin, Benin Republic. Hydrology. 2015;2(4):210-229.
- [9] Nagesh S, Laxmi D. Identifying suitable probability models for extreme flood heights in Mahanadi River Basin. Int. J. Agricult.Sci. 2021;17(1):197-207.
- [10] Nagesh S, Laxmi D.Estimating the extreme flood height quantiles using bayesian approach. Indian Journal of Ecology. 2023;50(1):210-214.
- [11] Provost SB, Saboor A, Cordeiro GM, Mansoor M.On the q-generalized extreme value distribution. REVSTAT-Statistical Journal. 2018;16(1):45-70.
- [12] Katz RW, Parlange MB, Naveau P.Statistics of extremes in hydrology. Advances in Water Resources. 2002;25(8-12):1287-1304.
- [13] HesarkazzaziS, ArabzadehR, Hajibabaei M, Rauch W, Kjeldsen TR, Prosdocimi I, Castellarin A, Sitzenfrei R.Stationary vs non-stationary modelling of flood frequency distribution across northwest England. Hydrological Sciences Journal. 2021;66(4):729-744.
- [14] Meis M, Llano MP, Rodriguez D.Quantifying and modelling the ENSO phenomenon and extreme discharge events relation in the La Plata Basin. Hydrological Sciences Journal. 2021;66(1):75-89.
- [15] Şen Z.Conceptual monthly trend polygon methodology and climate change assessments. Hydrological Sciences Journal. 2021;66(3):503-512.
- [16] Ferreira A, De Haan L.On the block maxima method in extreme value theory: PWM estimators. The Annals of statistics. 2015;43(1):276-298.
- [17] Dombry C.Maximum likelihood estimators for the extreme value index based on the block maxima method. arXiv preprint arXiv:1301.5611; 2013.
- [18] Yilmaz AG, Hossain I, Perera BJC. Effect of climate change and variability on extreme rainfall intensity– frequency–duration relationships: A case study of Melbourne. Hydrology and Earth System Sciences. 2014;18(10):4065-4076.
- [19] Coles S.Classical extreme value theory and models. In An introduction to statistical modeling of extreme values. Springer, London.2001;45-73.
- [20] Smith PJ. Maximum likelihood estimation of the distribution of length at age. Biometrics. 1987;601-615. __

© 2023 Nagesh and Dharmannavar; This is an Open Access article distributed under the terms of the Creative Commons Attribution License [\(http://creativecommons.org/licenses/by/4.0\)](http://creativecommons.org/licenses/by/3.0), which permits unrestricted use, distribution, and reproduction in any medium, provided the original work is properly cited.

Peer-review history: The peer review history for this paper can be accessed here (Please copy paste the total link in your browser address bar) https://www.sdiarticle5.com/review-history/110014